

# Generation of 100-year geomagnetically induced current scenarios

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## Abstract

100-year extreme geoelectric field and geomagnetically induced current (GIC) scenarios are explored by taking into account the key geophysical factors associated with the geomagnetic induction process. More specifically, we derive explicit geoelectric field temporal profiles as a function of ground conductivity structures and geomagnetic latitudes. We also demonstrate how the extreme geoelectric field scenarios can be mapped into GIC.

Generated statistics indicate 20 V/km and 5 V/km 100-year maximum 10-s geoelectric field amplitudes at high-latitude locations having poorly conducting and well-conducting ground structures, respectively. We show that geoelectric field magnitudes experience dramatic drop across boundary at about 40-60 degrees of geomagnetic latitude. We identify this as a threshold at about 50 degrees of geomagnetic latitude. The sub-threshold geoelectric field magnitudes are about an order of magnitude smaller than those at super-threshold geomagnetic latitudes.

The computed extreme GIC scenarios can be used in further engineering analyses that are needed to quantify the geomagnetic storm impact on the

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conductor systems such as high-voltage power transmission systems of interest. To facilitate further work on the topic, the digital data for generated geoelectric field scenarios are made publicly available.

## 1 Introduction

The potential for severe societal consequences has been driving recent increasing interest in extreme geomagnetic storm impact particularly on high-voltage power transmissions systems (e.g., National Research Council, 2008; North American Electric Reliability Corporation and the US Department of Energy, 2010). Although also GPS-based timing of the transmission system operations can be impacted by space weather storms, it is generally understood that geomagnetically induced currents (GIC) causing half-cycle saturation of high-voltage power transformer are the leading mode for the most severe problems such as electric blackouts and equipment damage (e.g., Kappenman, 1996; Molinski, 2002). Consequently, characterization of extreme GIC events is central for quantifying the technological impacts and societal consequences of extreme space weather events.

In this paper we investigate general characteristics of extreme geoelectric field and GIC events. Geoelectric field induced on the ground by spatiotemporally varying magnetospheric and ionospheric electric current systems is the primary physical quantity driving GIC and often a simple linear relationship is sufficient for mapping geoelectric field into GIC (e.g., Viljanen et al., 2006a; Pulkkinen et al., 2006; Ngwira et al., 2008; Pulkkinen et al., 2010). Consequently, the key challenge is to characterize extreme geoelectric field events, which is the primary goal of the paper.

Complete analysis of the risk from extreme space weather impacts on, for example, high-voltage power transmissions systems requires also engineering analyses specifying how given extreme GIC impacts the performance of the transformers and the system as a whole. Such a holistic definition of risk goes beyond the pure probability of occurrence of strong GIC events of certain magnitudes. It also comprises aspects of the vulnerability, i.e. how susceptible or robust today's power transmission systems behave toward the physical hazard of GIC. In addition, depending on parameters such as the geographical location or the time of day or the season, the impact of the same physically extreme (and hence rare) event might draw a very different picture. Such improved concepts are widely used in the natural catastrophe risk

modeling community (Grossi, 2005), e.g. in the insurance sector. However, today's natural catastrophe risk models have their main focus on more prevalent hazards, such as meteorological and hydrological events (Muir-Wood and Grossi, 2008) and geophysical events (e.g. earthquakes). Although more detailed engineering analyses are out of scope of the work at hand, our purpose is to facilitate further engineering and hazard analyses quantifying the risk extreme geomagnetic storms pose on high-voltage power transmission.

The geomagnetic induction process generating the ground geoelectric field is dependent both on the characteristics of geospace electric currents and on the local geological conditions dictating the electromagnetic response of the medium to geospace driving (for classic treatments see e.g., Wait, 1970; Berdichevsky and Zhdanov, 1984; Weaver, 1994). Consequently, the geoelectric field is a complex function of a number of geophysical factors that all need to be accounted for in the extreme event analysis. The key factors to consider are:

- The effect of the ground conductivity structure on the extreme geoelectric field amplitudes.
- The effect of the geomagnetic latitude on the extreme geoelectric field amplitudes.
- Temporal scales of the extreme events.
- Spatial scales of the extreme events.

Due to the lack of observational information about extreme events and due to the great variety of, for example, local geological conditions, accounting for all four factors above is a substantial challenge. Clearly, approximations and extrapolations are required in the analysis and the question thus becomes “what is the most feasible practical approach that also provides information directly usable in further engineering analyses?” These considerations lead us to use *extreme event scenario* approach. In this approach our goal is to generate several scenarios that represent the variability of the extreme events as a function of the four factors above. Further, the scenario approach will provide representative GIC time series that can be used directly in further engineering analyses.

Another question that needs to be addressed prior to our analyses is “what exactly is meant by an extreme event?” Extreme event can be defined

in a number of different ways and there has been discussion in the GIC community that one should look, for example, at “10 times March 13, 1989 event.” March 13, 1989 in this case refers to the space weather event that led to the collapse of the Hydro Quebec high-voltage power transmissions system in Canada (e.g., Bolduc, 2002). However, “10 times the March 13, 1989 event” is not a rigorous definition for an extreme event. First, one needs to define what physical parameters are used to amplify the March 13, 1989 event 10-fold. If one of the parameters is, for example, the *Dst* index which is the classic parameter used in measuring the geomagnetic storm strength, one is immediately faced with another problem. Namely, the minimum *Dst* index of the March 13, 1989 was -589 nT while the minimum *Dst* of the Carrington storm event of September 1-2, 1859 has been estimated to be approximately -850 nT (Siscoe et al., 2006). The Carrington geomagnetic storm event is the strongest in the recorded history and the minimum *Dst* of the event is only about 45 percent larger in absolute magnitude than *Dst* of the March 13, 1989 event. Consequently, amplifying the March 13, 1989 event 10-fold in terms of *Dst* index strength would quickly lead to unrealistic extreme storm scenarios. We note that similar argumentation applies also for the time derivative of the magnetic field often used as an indicator for GIC activity (e.g., Viljanen et al., 2001). First, there is no rigorous justification for arbitrarily amplifying the largest magnetic field fluctuations of the March 1989 event 10-fold. Second, despite the often good statistical association, the time derivative of the magnetic field is not the primary physical quantity driving GIC and consequently there is no direct one-to-one relation between the two parameters.

We will instead use a rigorous statistical definition for an extreme event and select physical parameter that is directly related to GIC. More specifically, we define an extreme event as maximum the 100-year amplitude of the 10 second resolution horizontal geoelectric field. The details of the definition will be discussed more in detail below but the basic philosophy for selecting a 100-year event is quite simple: we are looking for extreme events that occur significantly less frequently than once per solar cycle (i.e. 11 years) while at the same time being careful not to extrapolate the available observational information and statistics too far. As will be shown below, extrapolating the statistics to 100-year event is still reasonable while extracting information about significantly rarer events may not be feasible.

Section 2 of the paper describes the process used for generating the 100-year geoelectric field scenarios. We will first discuss the general philosophy

of the extreme scenario approach more in detail and then describe in Section 2.1 the generation of the baseline statistics. Subsequent sections describe the analyses associated with the four different factors: 2.2) the effect of the ground conductivity structure on the extreme amplitudes, 2.3) the effect of the geomagnetic latitude on the extreme amplitudes, 2.4) temporal scales of the extreme events and 2.5) spatial scales of the extreme events. Section 3 summarizes the generated extreme geoelectric field scenarios. In Section 4 we will describe how the geoelectric field scenarios can be mapped into GIC and Section 5 provides further discussion and outlines some of the work required for refining and improving the scenarios generated in this paper.

## 2 Generation of extreme geoelectric scenarios

As was explained above, we selected the extreme event scenario approach to account for the variability of geospace and geological conditions associated with extreme GIC events. More specifically, scenarios will be derived as a function of different representative ground conductivity structures and geomagnetic latitudes. Typical temporal scales of storm events will be captured by using temporal profiles from a representative storm event. 100-year scenarios are then achieved by scaling the representative storm event by the maximum amplitudes obtained via extrapolation of the geoelectric field amplitude statistics. In other words, we generate an artificial storm event that will produce the 100-year peak amplitudes in the statistics.

We chose to scale actual observed storm instead of using synthetic temporal profiles because of two major reasons. First, as will be discussed more in detail below, there are many types of dynamical processes in the solar wind-magnetosphere-ionosphere system that are capable of generating large GIC. All of these processes have their distinct spectral characteristics and consequently no single simple synthetic temporal profile is capable of capturing the full variability observed during extreme storms. The selected representative storm profiles instead include actual spectral signatures of many relevant geospace drivers of large GIC. Using actual observed storm events also captures the length of the extreme storm events, which may be important from the engineering analysis viewpoint.

It is important to note that linear scaling of representative storm event

carried out in this work is supported by the statistical results by Weigel and Baker (2003) and Pulkkinen et al. (2008). Weigel and Baker (2003) found that the shape of the probability distribution of high-latitude ground magnetic field fluctuations is nearly independent of solar wind driving conditions. Because the average solar wind state primarily enters through the standard deviation of the distributions derived by Weigel and Baker (2003), the solar wind input can be viewed as a linear amplifier of the high-latitude ground magnetic field fluctuations. In another words, one can select an arbitrary geomagnetic storm event and scale it to represent different solar wind driving conditions. The plane wave-based mapping, which has been shown to produce accurate modeled GIC, of the horizontal ground magnetic field components into the geoelectric field is a linear operation (see Appendix A). It follows that the results by Weigel and Baker (2003) hold also for the (modeled) geoelectric field.

Pulkkinen et al. (2008) on the other hand studied the statistics of modeled geoelectric field amplitudes at high-latitude locations. They found that the probability distribution of the geoelectric field amplitudes is approximately lognormal and that the shape of the distribution is nearly independent of both solar wind convective electric field amplitude and magnetospheric state measured in terms of the *Dst* index. Further, the mean of the distribution increased monotonically as the solar wind or magnetospheric conditions became more severe. This feature can be understood by considering lognormal probability distribution of variable  $x$

$$p \sim e^{-\frac{(\ln x - \mu)^2}{2\sigma^2}} \quad (1)$$

where  $\mu$  is the mean and  $\sigma^2$  the variance of the variable  $x$ 's natural logarithm, respectively. Linear amplification of the variable  $x$  by  $\alpha$  modifies the distribution as

$$p \sim e^{-\frac{(\ln x - \mu + \ln \alpha)^2}{2\sigma^2}} \quad (2)$$

In another words, the mean of the lognormal distribution is shifted. It follows that in terms of the shift in the mean of the lognormal distribution, the solar wind convective electric field and the *Dst* index can be viewed as a linear amplifier of the modeled high-latitude geoelectric field magnitudes. This finding is in a good agreement with the results by Weigel and Baker (2003) discussed above. Consequently, there is a good statistical justification for

linear scaling of the geoelectric field to represent the most extreme solar wind conditions responsible for the most extreme GIC.

## 2.1 Generation of the statistics

The basis of the statistical analysis is identical to that in Pulkkinen et al. (2008). We note that also, for example, Campbell (1980); Langlois et al. (1996); Boteler (2001) have studied the statistical aspects of extreme GIC events but these studies used much more limited datasets or relied on empirical relations to geomagnetic indices. In this work the statistics are generated by using 10-s geomagnetic field recordings from 23 high-latitude IMAGE magnetometer chain sites for the period of January 1993 to December 2006. The IMAGE stations are located in Northern Europe and cover about 55-75 degrees of geomagnetic latitude (corrected geomagnetic coordinates). Geomagnetic data from each IMAGE station are used to compute the local geoelectric field magnitudes  $E = |\mathbf{E}|$ , where  $\mathbf{E}$  is the horizontal geoelectric field. The horizontal geoelectric field is calculated by applying the plane wave method (see Appendix A). The plane wave method has been shown in numerous studies to be able to accurately map the observed ground magnetic field to the geoelectric field and observed GIC (e.g., Trichtchenko and Boteler, 2006; Viljanen et al., 2006a; Wik et al., 2008). Further, although the plane wave method assumes one-dimensional (1D) ground conductivity structure, the method has been shown to be applicable even in highly non-1D situations if an *effective* 1D ground conductivity is used (e.g., Thomson et al., 2005; Ngwira et al., 2008; Pulkkinen et al., 2010).

Pulkkinen et al. (2006) showed that while temporal averaging of the ground magnetic field from 1-s to 10-s has significant impact on peak time derivative of the ground magnetic field, the peak modeled geoelectric field amplitudes are not reduced significantly. Averaging the ground magnetic field below 10-s temporal resolution, however, was shown to impact the peak geoelectric field magnitudes (see Fig. 9 in Pulkkinen et al. (2006)). Consequently, 10-s data is used in the calculation of the geoelectric field. The ground conductivity structures discussed in the section below are used in the calculations and the results are used to generate the statistical occurrence of the modeled geoelectric field at the IMAGE stations.

## 2.2 The effect of the ground conductivity structure on the extreme amplitudes

The ground conductivity structure is one of the major factors impacting the geoelectric field magnitudes. The detailed electromagnetic response of the medium to geospace driving is dependent on the local ground conductivity structure and consequently accurate calculation of the local geoelectric field and GIC requires knowledge about the local geological conditions. Consequently, we follow the approach by Pulkkinen et al. (2008) and select two ground conductivity models representing realistic extreme ends of conducting (British Columbia, Canada) and resistive (Quebec, Canada) grounds. The resistive Quebec model, which is associated with larger geoelectric field amplitudes, will be associated with a scenario having the most extreme GIC.

It is noted that, strictly speaking, one cannot apply Canadian ground models to geomagnetic observations from an entirely different geographical region as was done by Pulkkinen et al. (2008) and as is done here. However, to a good approximation the same magnetospheric-ionospheric source current will produce similar total magnetic field variations at regions with different ground conductivity structures. Consequently, a deviation from the strictly consistent approach in using the ground models and geomagnetic field observations is justified. See Pulkkinen et al. (2008) for a more detailed discussion on this.

Fig. 1 shows the statistical occurrence of the geoelectric field at IMAGE stations for the two ground conductivity structures. Approximate visual extrapolations to 100-year peak magnitudes are also given. As is seen from Fig. 1, the peak magnitudes for different stations group quite tightly and extracting 100-year values requires extrapolation over about one order of magnitude in occurrence rates. Further, the tails of the occurrence distributions fall off with fairly continuous slope and it is thus argued that while extraction of much rarer events may be a challenge, the presented extrapolations to 100-year occurrence rates is reasonable. While one could try applying a more rigorous extreme value theory (EVT) for studying the tails of the occurrence distributions, we argue that given the approximate nature of the work at hand, visual extrapolation is perfectly sufficient; it is unlikely that any reasonable EVT fitting procedure would give extreme amplitudes out of the upper and lower limits indicated in Fig. 1.

As is seen from Fig. 1, for poorly conducting (represented by Quebec ground model) high-latitude regions the maximum 100-year amplitude of the



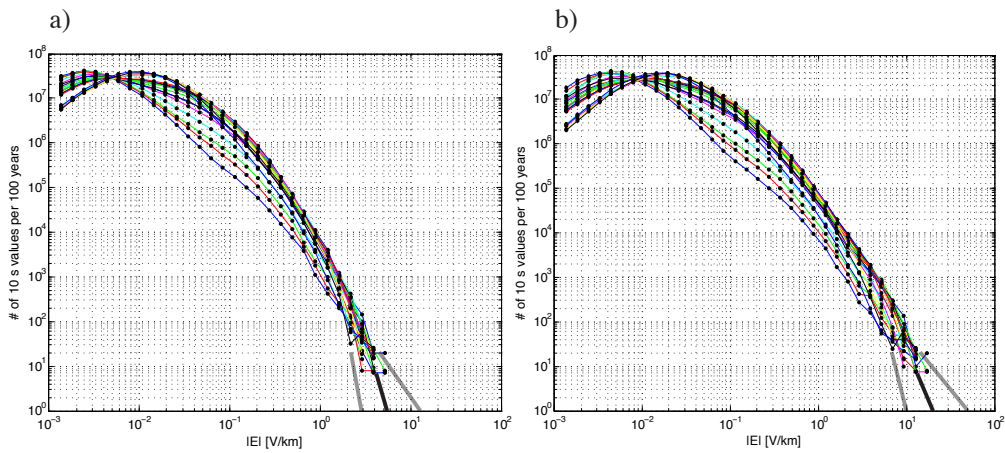


Figure 1: Statistical occurrence of the geoelectric field computed using the ground conductivity structure of a) British Columbia, Canada and b) Quebec, Canada. Different curves correspond to different IMAGE stations used in the computation of the geoelectric field. The thick black lines indicate approximate visual extrapolations of the statistics to 100-year peak magnitudes. The thick grey lines indicate the reasonable lower and upper boundaries for the extrapolated values. The figure is modified version of Fig. 2 in Pulkkinen et al. (2008).

10 second resolution horizontal geoelectric field is estimated to be between 10-50 V/km. For well-conducting regions (represented by British Columbia ground model) the maximum amplitudes are about factor of 5 smaller and estimated to be between 3-15 V/km.

### 2.3 The effect of the geomagnetic latitude on the extreme amplitudes

Due to the location of the IMAGE magnetometer stations, the analysis in Section 2.2 applies directly only to high-latitude locations between 55-75 degrees of geomagnetic latitude. Different magnetosphere-ionosphere source currents dominate the ground magnetic field signature at different geomagnetic latitudes. For example, at high-latitudes magnetic signature is dominated by auroral ionospheric currents while at low-latitudes the signature is combination of multiple sources such as ring, magnetopause, magnetotail and the equatorial electrojet currents (e.g., Ohtani et al., 2000). Further, different magnetosphere-ionosphere current systems have their own spatiotemporal characteristics and consequently it is central to account for the geospace source variability in the generation of the extreme geoelectric field amplitudes.

Unfortunately, global 10-s ground magnetic field observations are not available for extended time periods. The standard temporal resolution, for example, for the International Real-time Magnetic Observatory Network (INTERMAGNET) sites is 60-s. Consequently, global investigations of extreme geoelectric field amplitudes are confined to use 60-s resolution data at best, which may cut some of the peak geomagnetic field fluctuation and geoelectric field amplitudes. Consequently, we assume that the *relative* change of the peak amplitudes as a function of geomagnetic latitude is the same for 10-s and 60-s data. While validity of this assumption cannot be verified easily for the time derivative of the magnetic field, since the geoelectric field is not as sensitive to temporal averaging (Pulkkinen et al., 2006), we argue that the available 60-s temporal resolution is sufficient for the purpose of this part of the work.

We studied the global behavior of the ground magnetic field and geoelectric field fluctuations for two extreme geomagnetic storm events of special significance: March 13-15, 1989 and October 29-31, 2003. The March 1989 storm caused the collapse of the Hydro Quebec high-voltage power trans-

mission system while the October 2003 storm caused the blackout in Southern Sweden (Pulkkinen et al., 2005) and possibly problems with the South African high-voltage transmission system (Gaunt and Coetzee, 2007). Both storms were generated by major solar coronal mass ejection events known to be the most significant driver of large GIC (e.g., Borovsky and Denton, 2006; Kataoka and Pulkkinen, 2008; Huttunen et al., 2008). The minimum  $Dst$  indices of March 1989 and October 2003 storms were -589 nT and -383 nT, respectively. Using the  $Dst$  index as a measure of the storm strength, March 1989 and October 2003 storms rank between years 1957-2010 for which  $Dst$  data is available as 1st and 8th strongest, respectively. In fact, as can be seen from Fig. 2 showing the statistical occurrence of hourly  $Dst$  values between 1957-2010, the peak  $Dst$  of the March 1989 storm may have been close to the 100-year amplitude. Tsubouchi and Omura (2007) used extreme value statistics to estimate that the  $Dst$  of the storm was 60-year event. Further, the statistics suggest that the Carrington 1859 storm peak  $Dst$  of about -850 nT could in fact be rarer than a 100-year event. It is, however, noted that the transition in the slope of the distribution in Fig. 2 at about 300 nT indicates that finite size of the sample may hinder the accurate estimation of the characteristics of the tail of the distribution. Consequently, one should be careful in making interpretations about the likelihoods of the extreme  $Dst$  values.

We retrieved 60-s global geomagnetic field data from INTERMAGNET ([www.intermagnet.org](http://www.intermagnet.org)) for the two months containing the two storm events and removed a visually determined baseline from the observations. We checked the data for clear bad values, and stations with suspicious data were removed from the analysis. Short data gaps were patched using linear interpolation. The Quebec ground conductivity model was then applied with the plane wave method to compute the geoelectric field at each station.

Fig. 3 shows the spatial distributions of the maximum computed geoelectric field and the maximum time derivative of the horizontal magnetic field taken over the March 13-15, 1989 and October 29-31, 2003 events. Fig. 4 in turn shows the latitude distributions of the maximum geoelectric field, the maximum time derivative of the horizontal magnetic field and the maximum amplitude of the horizontal magnetic field. As can be seen from Fig. 3, the global coverage of the magnetometer stations for both events is fairly good. Quite interestingly, both Figs. 3 and 4 indicate dramatic global drop in the maximum magnitudes of all three parameters approximately between 40-60 degrees of geomagnetic latitude. The maximum geoelectric field, the

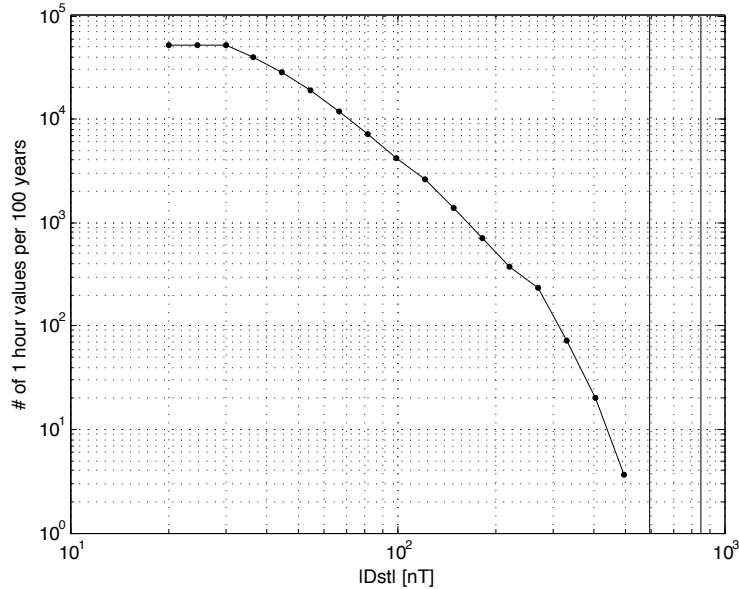


Figure 2: Statistical occurrence of hourly  $Dst$  index values between years 1957-2010. The two vertical lines indicate  $|Dst|$  of 589 nT and 850 nT.

maximum time derivative of the horizontal magnetic field and the maximum amplitude of the horizontal magnetic field all experience approximately an order of magnitude drop across the threshold at about 50 degrees of geomagnetic latitude. Furthermore, the same drop is observed both in the northern and southern hemispheres.

Figs. 3b, 3b, 4b and 4d show confined enhancement of maximum computed geoelectric field and the maximum time derivative of the horizontal magnetic field for two stations at about magnetic equator. The maximum values occurred between 08-12 magnetic local time, which indicates that the enhancement may be associated with equatorial electrojet that is a localized band of ionospheric current between about -5 to 5 degrees of geomagnetic latitude (Lühr et al., 2004). This indicates, to our knowledge for the first time, that also equatorial electrojet is capable of generating notable GIC. However, more detailed study out of scope of the work in this paper is required to confirm if equatorial electrojet indeed drives enhanced GIC at the magnetic equator.

One of the interesting features in Figs. 3 and 4 is the implied universality of the threshold at about 50 degrees of geomagnetic latitude. In terms of

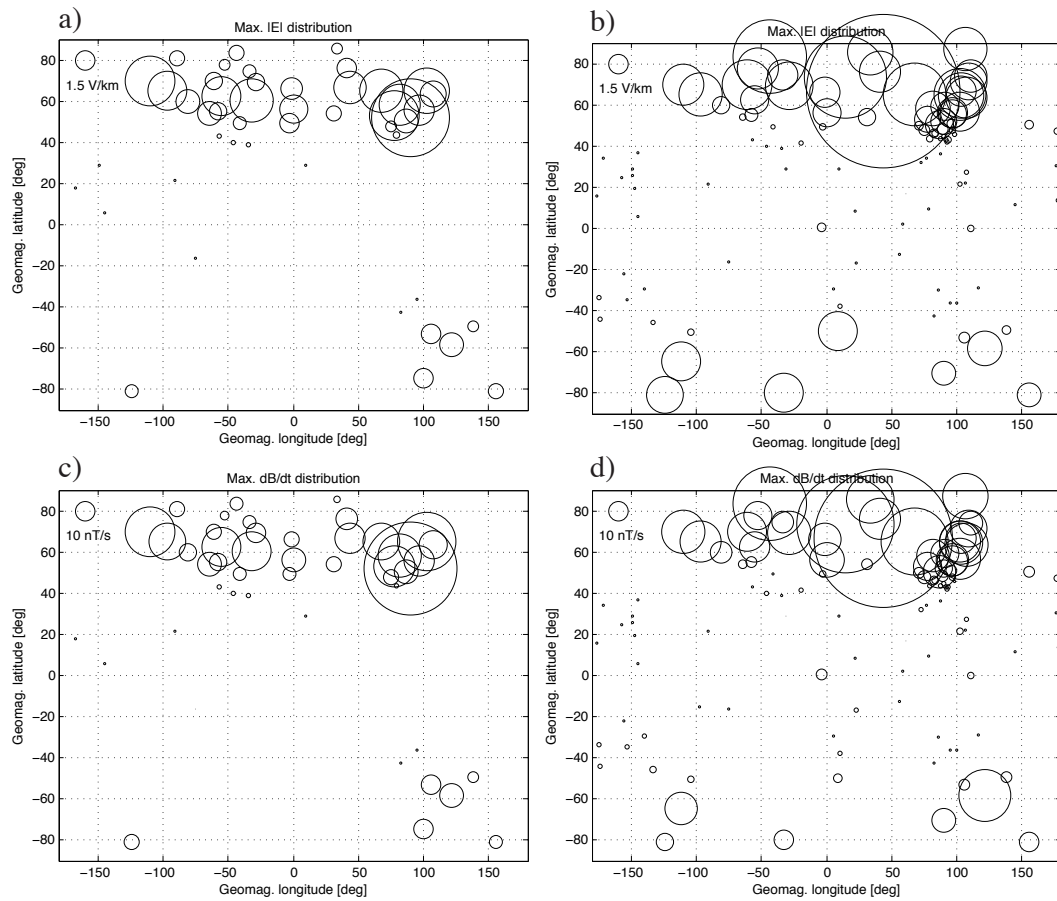


Figure 3: Spatial distributions of the maximum computed geoelectric field (top row) and the maximum time derivative of the horizontal magnetic field (bottom row) for March 13-15, 1989 (panels a and c) and October 29-31, 2003 (panels b and d) events. The center of each circle indicates the location of the corresponding magnetometer station and the radius of the circle indicates the maximum magnitudes of the physical parameters.

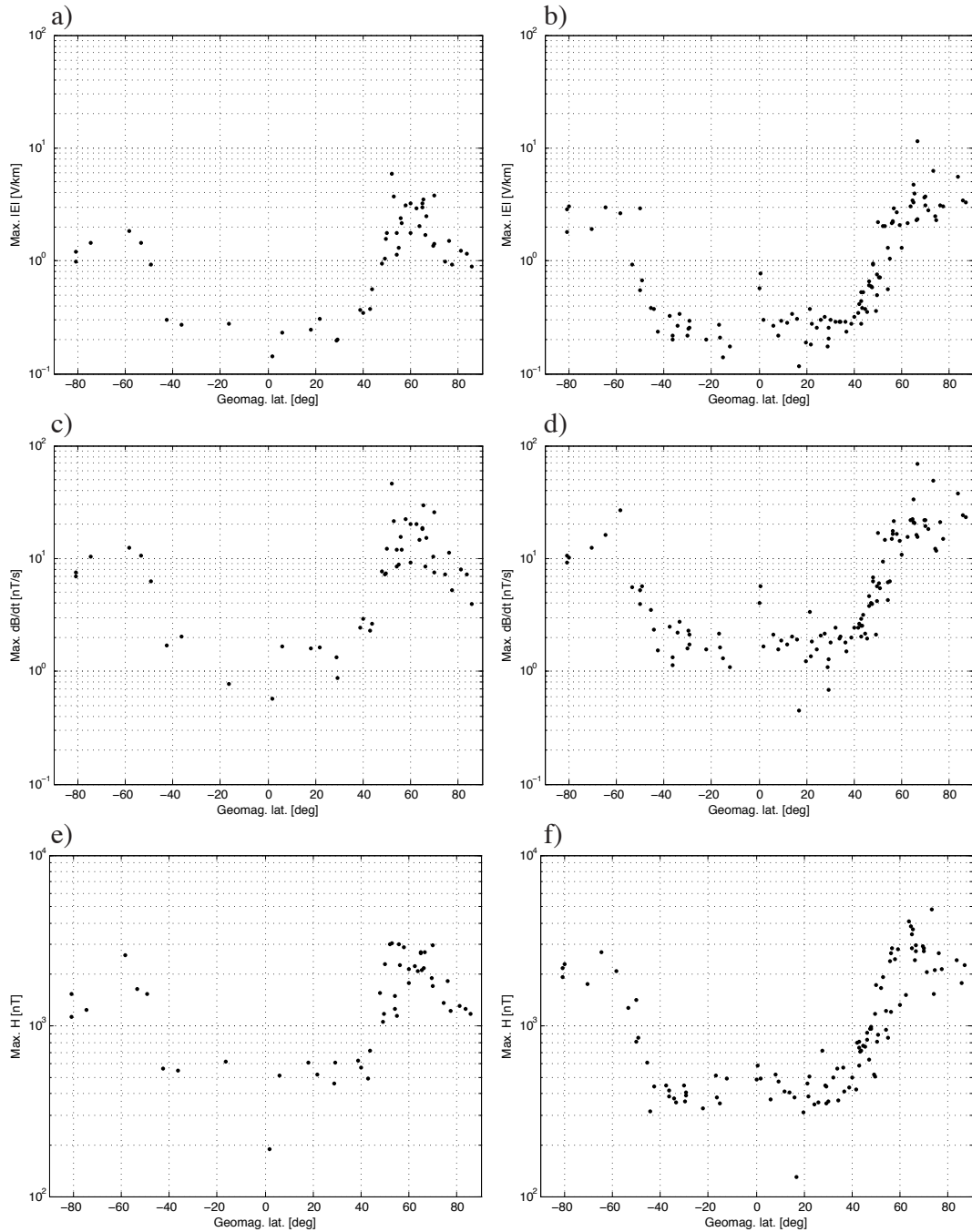


Figure 4: Geomagnetic latitude distributions of the maximum computed geoelectric field (top row), the maximum time derivative of the horizontal magnetic field (middle row) and the maximum amplitude of the horizontal magnetic field (bottom row) for March 13-15, 1989 (panels a, c and e) and October 29-31, 2003 (panels b, d and f) events.

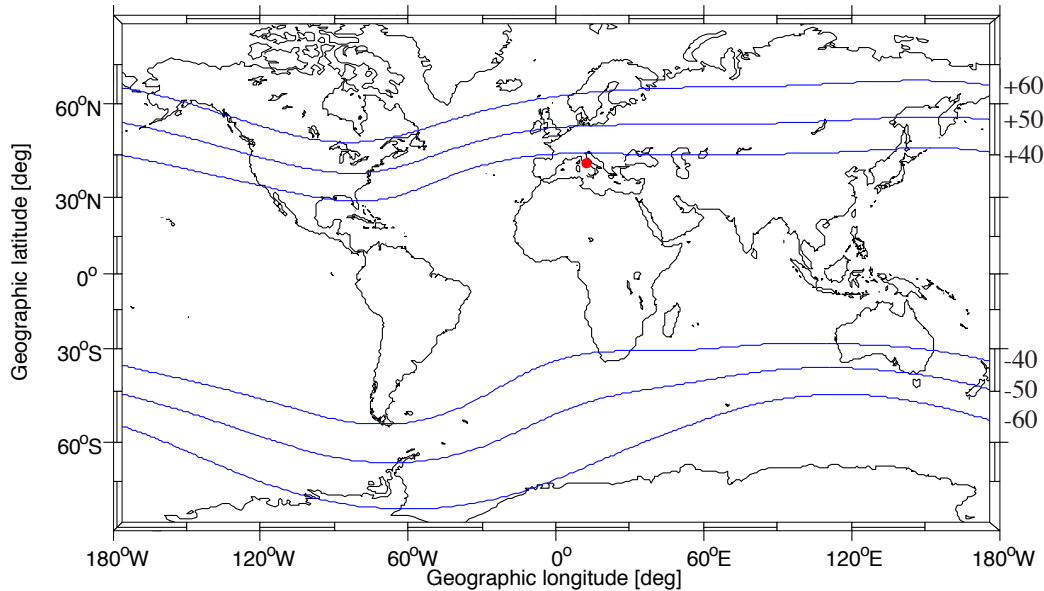


Figure 5: -60, -50, -40, 40, 50 and 60 degrees of geomagnetic latitude contours (blue lines) displayed on geographic map. The red dot indicates the location of Rome, Italy. Altitude Adjusted Corrected Geomagnetic Coordinate coefficients for year 2000 were used in the generation of the contours.

*Dst* index the March 1989 storm was significantly stronger than the October 2003 storm. Consequently, one could expect that, for example, the auroral boundaries would have penetrated to significantly lower latitudes during the March 1989 event. However, the drop of the maximum amplitudes occurs for both storm events approximately between 40-60 degrees of geomagnetic latitude. Clearly usage of global geomagnetic data for extended time periods is required to study and confirm the location and dynamics of the threshold geomagnetic latitude more definitively. However, Thomson et al. (2011) used 28 years of 60-s geomagnetic data from Europe and found a similar threshold at about 55 degrees of geomagnetic latitude for the time derivative of the horizontal magnetic field. We thus conjecture that the threshold at about 50-55 degrees of geomagnetic latitude is a universal feature of most major or extreme geomagnetic storms (for geographical reference, see Fig. 5).

As is seen from Fig. 4, the threshold latitude for the maximum geoelectric field and the maximum time derivative of the horizontal magnetic field is reflected also in the maximum amplitude of the horizontal magnetic field.

This is an indication that the most extreme magnetic field fluctuations and geoelectric field magnitudes are associated with the auroral current system that is known to be responsible for the largest perturbations of the ground magnetic field. For example, while magnetospheric ring current can generate horizontal magnetic field perturbations of the order of hundreds of nT, auroral currents regularly generate perturbations of the order thousands of nT. Consequently, the question about the location of the threshold geomagnetic latitude can be cast also in terms of maximum possible expansion of the auroral current system.

Possibly the strongest geomagnetic storm in the recorded history is the Carrington event of September 1-2, 1859 (e.g., Tsurutani et al., 2003; Siscoe et al., 2006). The minimum estimated  $Dst$  index of the storm was -850 nT and there were (poleward horizon) auroral sightings from as low as 23 degrees of geomagnetic latitude. However, from the viewpoint of the analysis in this paper, perhaps the most significant observation during the event comes from Rome, Italy (see Fig. 5). More specifically, perturbation of the order of 3000 nT was observed using a bifilar magnetometer that indicates relative changes in the horizontal magnetic field strength (Loomis, 1860). The geomagnetic latitude of Rome, Italy is about 36 degrees and comparing this to Figs. 4e and 4f, the observation indicates that the maximum expansion of the auroral current system may have been about 20 degrees more southward than during the March 1989 or October 2003 storms. Although this may sound somewhat fantastic and the single data point was not based on modern scientific instrumentation, one cannot simply disregard the Rome observation. It may thus be possible that during the most extreme geomagnetic storms the auroral current system and the accompanying extreme geoelectric fields and GIC can penetrate significantly below the threshold at about 50 degrees of geomagnetic latitude discussed above. Unfortunately, due to the poor spatial coverage, low temporal sampling rates and off-scale magnitudes, magnetic recordings of the Carrington event do not allow for more detailed analysis of the global geoelectric field and GIC characteristics (Nevanlinna, 2006; Boteler, 2006; Nevanlinna, 2008). At sub-auroral latitudes in Finland and Russia, the greatest measured hourly point deviations in the horizontal field during the event were about 1000 nT (Nevanlinna, 2008).

Observations of low-latitude boundary of the auroral emissions indicative of the general location of the auroral region provide an alternative view to question about of the threshold geomagnetic latitude. Records of auroral sightings are available also for historical storms and these enable approximate



reconstructions of the auroral region morphology during the corresponding storms. For example, Silverman and Cliver (2001); Cliver and Svalgaard (2004); Silverman (2006) provide auroral data collected from numerous catalogues and earlier studies. Most importantly, Silverman and Cliver (2001) provide maximum equatorward extent of the visual aurora for three historical extreme geomagnetic storms: August 28-29, 1859, September 1-2, 1859 and May 14-15, 1921. While August 28-29, 1859 event was part of the August-September 1859 extreme storm sequence, analysis by Kappenman (2006) indicated that minimum  $Dst$  index of the May 14-15, 1921 event may have been comparable to that of the Carrington event. Silverman and Cliver (2001) provide auroral boundary locations both for the overhead auroras and auroras observed in the poleward horizon. However, only the overhead auroral sightings provide unambiguous determination of the auroral boundary location and thus auroras observed in the poleward horizon are not discussed further here.

Table 1 shows the approximate low-latitude auroral boundary locations for the four extreme storm event. The boundary location for the March 13-15, 1989 was determined by visual inspection of Dynamic Explorer 1 (DE-1) ultraviolet auroral emission imaging data in Allen et al. (1989) for March 14, 1989 01:51 UT. The visually determined boundary at about 40 degrees of geomagnetic latitude is in a very good agreement with the electron precipitation boundary determined from low-Earth orbit Defense Meteorological Satellite Program (DMSP) satellite data by Yokoyama et al. (1998).

Table 1: The approximate maximum equatorward auroral boundary locations of selected extreme geomagnetic storm events. The table is based on data in Allen et al. (1989); Yokoyama et al. (1998); Silverman and Cliver (2001). <sup>a</sup>Maximum equatorward extent of the overhead visual aurora. <sup>b</sup>Maximum equatorward extent in DE-1 imagery and electron precipitation boundary from DMSP.

Event Date	Location in geomagnetic latitude
August 28-29, 1859	48 deg <sup>a</sup>
September 1-2, 1859	41 deg <sup>a</sup>
May 14-15, 1921	40 deg <sup>a</sup>
March 13-15, 1989	40 deg <sup>b</sup>

The striking feature of the data in Table 1 is that the boundary is confined

to approximately 40 degrees of geomagnetic latitude or greater for all four events. This is also the approximate low-latitude boundary for the transition from low-latitude maximum field magnitudes to high-latitude maximum field magnitudes seen in Fig. 4. As auroral emissions are congruent with the auroral ionospheric current fluctuations, low-latitude auroral boundary locations for extreme storm events in Table 1 along with “calibration” to modern magnetic field observations via March 13-15 1989 event in Fig. 4 provide further indication of possible generality of the threshold at about 50 degrees of geomagnetic latitude. Although their results were based on data only for year 1998, Ahn et al. (2005) also concluded that the lowest possible latitude of the center of the ionospheric westward electrojet seems to be at around 60 degrees of geomagnetic latitude, which is consistent with Figs. 3 and 4. Ahn et al. (2005) also showed that the low-latitude boundary of auroral emissions tends to locate equatorward of the westward electrojet, which is also consistent with our findings above.

It is noted that the Rome observation of geomagnetic field perturbation of the order of 3000 nT during the Carrington event contradicts with the idea of the threshold geomagnetic latitude at about 50 degrees given the 40 degrees of geomagnetic latitude boundary for the event in Table 1. The possible explanations for the discrepancy are 1) low-latitude boundary of the auroral emissions is not always congruent with the low-latitude boundary of the auroral ionospheric currents, 2) the actual low-latitude boundary of the overhead auroral emissions was at the time of the Rome observation lower than 40 degrees of geomagnetic latitude but was not captured by any of the historical records and 3) the Rome magnetic field observation was erroneous. Although especially possibility 1) seems unlikely, further work is needed to find the most plausible explanation for the discrepancy.

Although the Rome geomagnetic field recordings for the Carrington event indicate that it may be possible for the auroral current system to penetrate significantly lower than to about 50 degrees of the geomagnetic latitude, there is no direct means to quantify the likelihood of such occurrence. Further, the approximate low-latitude auroral boundary locations for some historical extreme events provide indication of maximum expansion of the auroral region to about 40 degrees of geomagnetic latitude, which is consistent with the view in Figs. 3 and 4. Consequently, given the very similar properties of the March 1989 and October 2003 storms in Figs. 3 and 4 along with statistics by Thomson et al. (2011), we are inclined to hold our conjecture that the threshold at about 50-55 degrees of geomagnetic latitude holds for most

major and extreme geomagnetic storms - possibly also for 100-year events. In terms of extreme event scenarios and scaling this means that the extreme geoelectric field amplitudes experience about factor of 10 drop across the region from 60 to 40 degrees of geomagnetic latitude (see Fig. 5). However, it is emphasized again that more extreme geomagnetic storm data is required for more definite conclusions regarding the 100-year location of the threshold geomagnetic latitude.

## 2.4 Temporal scales of the extreme events

Many different types of dynamical processes in the solar wind-magnetosphere-ionosphere system are capable of generating large GIC. For example, auroral substorms, geomagnetic pulsations, sudden impulses and enhancements of magnetospheric convection-related auroral electrojets are known to drive GIC (e.g., Kappenman, 2003; Pulkkinen et al., 2003; Viljanen et al., 2006b). All of these processes have their distinct spectral characteristics and thus no single simple synthetic temporal profile is capable of capturing the full temporal variability observed during extreme storms (e.g., Pulkkinen and Kataoka, 2006; Kataoka and Pulkkinen, 2008). Consequently, we chose to use actual geomagnetic storm event to provide a representative temporal profile for the generated scenarios. We selected 10-s geomagnetic field observations from Nurmijärvi Geophysical Observatory, Finland and Memanbetsu Geophysical Observatory, Japan for the period of October 29-31, 2003 to provide the temporal profiles for the scenarios. The Nurmijärvi Geophysical Observatory is located approximately at 57 degrees of geomagnetic latitude and was thus within the region experiencing the most extreme magnetic field fluctuations above the threshold geomagnetic latitude during the storm. Memanbetsu Geophysical Observatory in turn is located approximately at 35 degrees of geomagnetic latitude and was thus within the region experiencing the magnetic field fluctuations below the threshold geomagnetic latitude during the storm. We also confirmed that the selected representative storm profiles include signatures of auroral substorms, geomagnetic pulsations, sudden impulses and enhancements of magnetospheric convection-related auroral electrojets. The selected magnetometer stations and the geomagnetic storm event thus provide a good representation of the ground electromagnetic field fluctuations during major geomagnetic storms.

Geomagnetic field observations from Nurmijärvi and Memanbetsu Geophysical Observatories for October 29-31, 2003 were applied with the Que-

bec ground conductivity model and the plane wave method to map the geomagnetic field into the horizontal geoelectric field. The obtained geoelectric field time series was then normalized so that the maximum amplitude of the signal is exactly 1 (Fig. 6), i.e.

$$\max(\sqrt{E_x^2 + E_y^2}) = 1 \quad (3)$$

where  $E_x$  and  $E_y$  are the normalized horizontal geoelectric field components and the maximum is taken over the storm event. The normalized horizontal geoelectric field is the signal that is used to scale to different maximum 100-year amplitude scenarios as a function of ground conductivity structures and geomagnetic latitudes.

## 2.5 Spatial scales of the extreme events

In the challenge of generating 100-year geoelectric field and GIC scenarios, characterizing the spatial scales of the extreme events may be the most difficult task. As was discussed above, many different types of processes in the solar wind-magnetosphere-ionosphere system drive large GIC and each of these processes have their characteristic temporal and spatial scales. Especially the global spatial scales of the geoelectric field associated with these processes are generally not well-known. However, high-latitude (auroral) magnetic field and geoelectric field fluctuations tend to be often poorly correlated over distances larger than of the order of 100 km (Pulkkinen et al., 2007, and references therein). Since on the other hand the extreme geomagnetic field and geoelectric field fluctuations are associated with the enhancements of the auroral current system that can be global, these two aspects give rise to a two-fold view: while large or extreme geoelectric field magnitudes can be experienced across the globe in the region covered by the auroral current system, the spatial correlation lengths associated with the field fluctuations can be short. Further complications are caused by the horizontal variations in the ground conductivity structure. The electromagnetic response to geospace driving is a strong function of the ground conductivity structure and steep horizontal conductivity gradients can generate steep fluctuations in the spatial geoelectric field structure.

Since we have no means of generating a global geoelectric field structure that would represent with any reasonable accuracy the true spatial scales (and spatial correlations) of the extreme fields, we are constrained to rep-

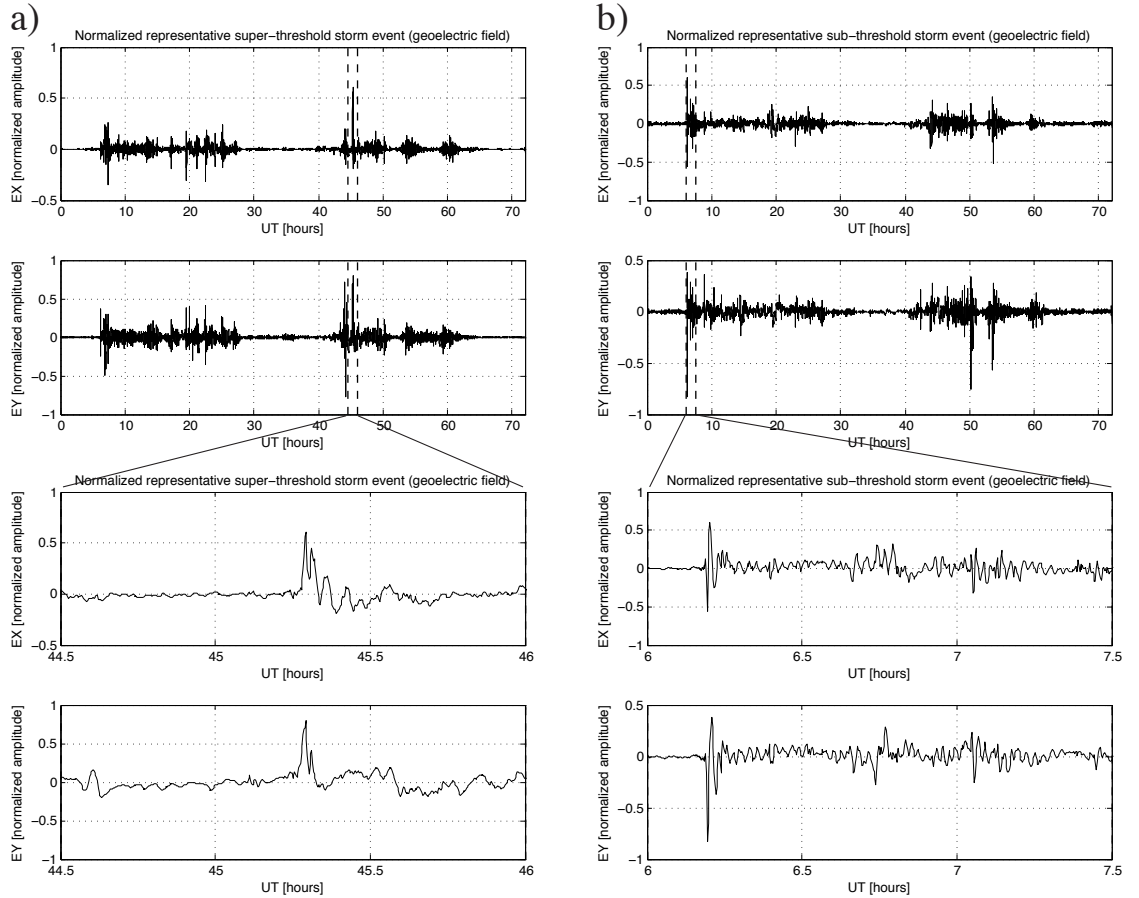


Figure 6: Normalized representative horizontal geoelectric field components ( $X$  indicates geographic north,  $Y$  indicates geographic east) for the full storm event and 1.5 hour long subsection containing the maximum field magnitude for super-threshold geomagnetic latitude locations represented by Nurmijärvi Geophysical Observatory (panel a) and sub-threshold geomagnetic latitude locations represented by Memanbetsu Geophysical Observatory (panel b). The time is hours from October 29, 2003 00:00 UT. The maximum geoelectric field magnitude at Nurmijärvi was caused by an auroral substorm while the maximum field magnitude at Memanbetsu was caused by a sudden impulse. See the text for details.

resent fields in regional scales. Consequently, we will assume that the geoelectric field is uniform over spatial scales of the order of 100-1000 km. In another words, the geoelectric field of the extreme storm scenarios has the same instantaneous direction and magnitude throughout the region of interest. The spatial uniformity amounts to assuming that there are no significant horizontal variations in the ground conductivity structure or in the spatial structure of the source field fluctuations. It is emphasized that while these assumptions may be reasonable for the purpose of the extreme scenarios on regional scales, they will break and should not be used in global scales.

### **3 Summary of the extreme geoelectric field scenarios**

Summarizing the findings in Section 2, the four key factors introduced in Section 1 are addressed in the extreme geoelectric field scenarios the following ways:

- The effect of the ground conductivity structure on the extreme geoelectric field amplitudes - two ground conductivity models representing realistic extreme ends of conducting and resistive grounds were applied with the IMAGE magnetometer data and the plane wave method. The results were used to estimate 100-year amplitudes of the 10 second resolution horizontal geoelectric field at high-latitudes. The resistive ground model is associated with approximately factor of five larger geoelectric field amplitudes than the conducting ground model.
- The effect of the geomagnetic latitude on the extreme geoelectric field amplitudes - we identified a threshold geomagnetic latitude across which the maximum geoelectric field amplitudes experience approximately an order of magnitude decrease.
- Temporal scales of the extreme events - representative time series from selected magnetometer stations for a major event storm event were used to provide realistic temporal profiles. Stations above and below the identified threshold geomagnetic latitude are used.
- Spatial scales of the extreme events - we assume spatially uniform geoelectric field structure in regional scales.

Fig. 7 summarizes the generated four 100-year geoelectric field scenarios. In Fig. 7a and Fig. 7b the normalized geoelectric field in Fig. 6a was scaled using the maximum amplitudes of 20 V/km and 5 V/km obtained from the high-latitude statistics in Fig. 1. In Fig. 7c and Fig. 7d the normalized geoelectric field in Fig. 6b was scaled by order of magnitude smaller maximum field strengths for sub-threshold geomagnetic latitudes. The threshold geomagnetic latitude can be set to 50 degrees or for more conservative estimates to 40 degrees of geomagnetic latitude. The geoelectric field is assumed spatially uniform in regional scales for all scenarios in Fig. 7.

## 4 Mapping geoelectric field scenario to geomagnetically induced currents

Mapping the extreme geoelectric field scenarios into GIC is a highly system-dependent operation. The response of the conductor system is dependent on the electrical characteristics and topology of the system and consequently it is generally speaking not feasible to provide any “prototype” configuration that could be applied to a variety of different situations. In another words, one needs to have additional engineering information available about the characteristics of the conductor system of interest prior to mapping geoelectric field into GIC.

Once the engineering information about the system has been acquired, there are two fairly straightforward means to carry out the mapping. First, if computation of GIC distribution throughout the (regional) system is needed, one can apply techniques by Lehtinen and Pirjola (1985) for discretely grounded systems such as high voltage transmission systems and so-called distributed source transmission line (DSTL) theory for continuously grounded systems such as buried oil and gas pipelines (Boteler, 1997).

To demonstrate the application of the extreme geoelectric field scenarios in computing GIC distribution throughout a high-voltage power transmission system, we considered Dominion’s Virginia Power grid model shown in Fig. 8. The model is built based on a DC-mapping of Dominion’s high-voltage transmission network. Typically, due to the scales associated with the GIC phenomenon, DC models should contemplate the 500 kV and the 230 kV networks; a few key 115 kV transmission lines are also considered in this

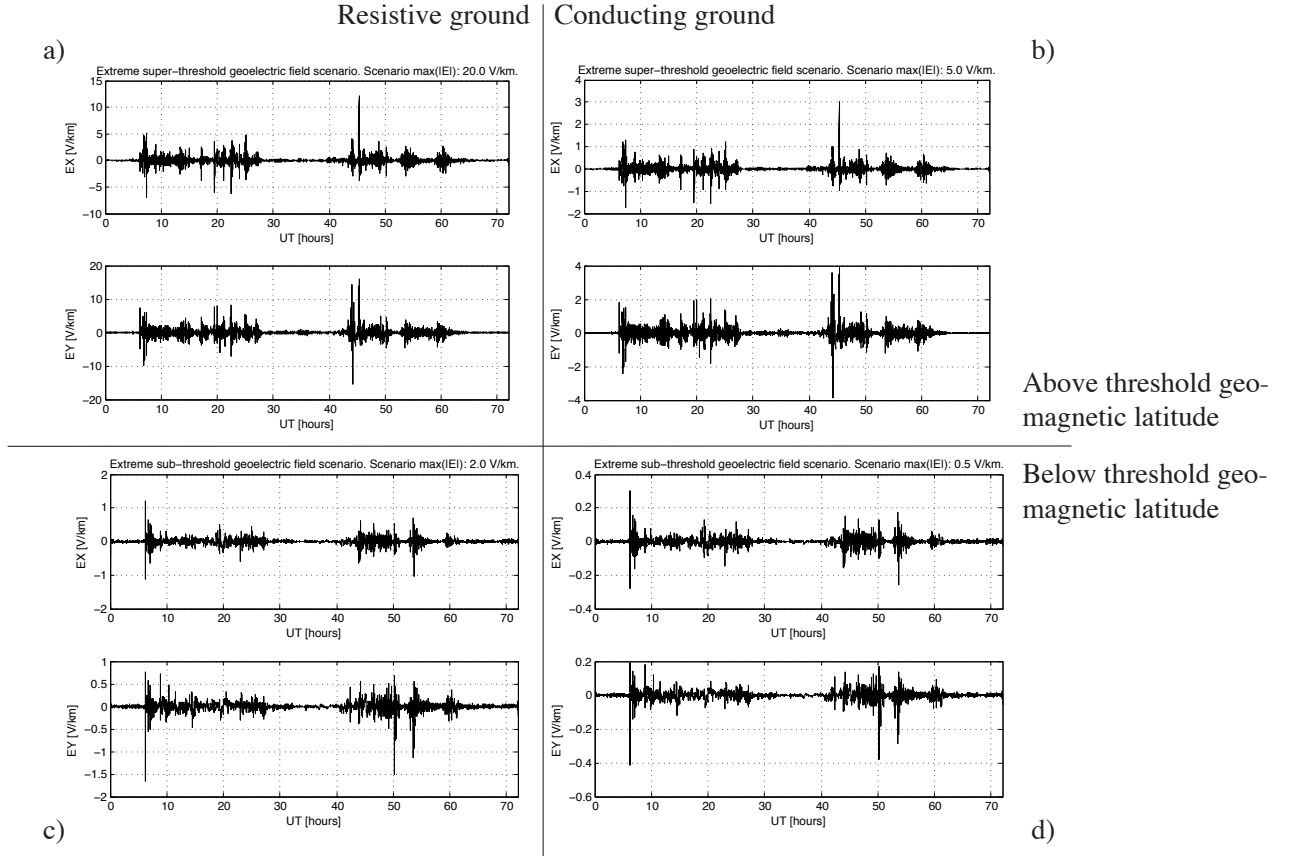


Figure 7: Illustration of extreme horizontal geoelectric field scenarios ( $X$  indicates geographic north,  $Y$  indicates geographic east). Panel a): Scenario for resistive ground structures for locations above the threshold geomagnetic latitude. The maximum geoelectric field amplitude is 20 V/km. Panel b): Scenario for conductive ground structures for locations above the threshold geomagnetic latitude. The maximum geoelectric field amplitude is 5 V/km. Panel c): Scenario for resistive ground structures for locations below the threshold geomagnetic latitude. The maximum geoelectric field amplitude is 2 V/km. Panel d): Scenario for conductive ground structures for locations below the threshold geomagnetic latitude. The maximum geoelectric field amplitude is 0.5 V/km. Note that the vertical scales are different in different panels.



particular model. Since the geoelectric field is assumed spatially uniform across the system, transmission lines between substations can be represented as straight lines. At each system inter-tie, a system equivalent is modeled as a low resistance to the ground ( $0.1 \Omega$ ).

The calculation of GIC flows in Dominion’s high-voltage transmission system is performed by using the matrix formulation derived in Lehtinen and Pirjola (1985). In general, GIC flows are a function of system topology, line resistances and geospatial orientation, transformer type and winding resistance, grounding resistance, series line compensation, and of course, geoelectric field. Fig. 8 shows a snapshot of GIC flows at each substation, in a per-phase basis, caused by the maximum amplitude of the geoelectric field in the scenario shown in Fig. 7c. The selection of the storm scenario is based on Dominion’s geomagnetic latitude (below the threshold latitude) and ground conductivity (resistive ground). To demonstrate that the approach provides a time series of GIC throughout the system over the entire storm scenario, Fig. 9 shows times series of modeled GIC flow in one of Dominion’s transformers. Note that the approach provides corresponding time series for any location in the high-voltage transmission system.

In an alternative approach, if only local GIC flowing through, for example, individual node of power transmission system is needed, one can apply simple linear relation

$$GIC = aE_x + bE_y \quad (4)$$

where  $(E_x, E_y)$  are the horizontal components of the geoelectric field and  $(a, b)$  the system parameters.  $(a, b)$  that depend on the topology and electrical characteristics of the conductor system under investigation can be derived for individual locations by using information about the full conductor system (Pulkkinen et al., 2006) or by inverting the parameters from GIC and ground magnetic field observations (Pulkkinen et al., 2007). The linear relation in Eq. (4) has been shown in numerous studies to hold to a good approximation in many situations of interest (e.g., Pulkkinen et al., 2006; Ngwira et al., 2008; Pulkkinen et al., 2010). The typical values for  $(a, b)$  are in the range between 0-200 A·km/V (Pulkkinen et al., 2008, and references therein). For example, using mid-range  $a = b = 50$  A·km/V one gets the extreme GIC scenario in Fig. 10 for conductive ground at location above the threshold geomagnetic latitude.

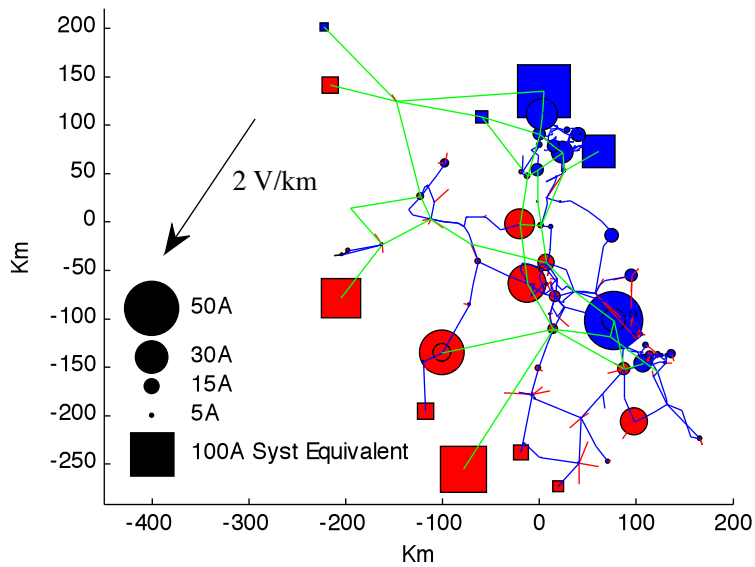


Figure 8: Modeled GIC distribution in the Dominion's Virginia Power high-voltage transmission system for hour 6.17 in the scenario shown in Fig. 7c. Green, blue and red lines indicate 500 kV, 230 kV and 115 kV transmission lines, respectively. The back arrow indicates the direction and magnitude of the horizontal geoelectric field and blue and red circles indicate the magnitude of GIC flowing from the ground to the grid and from the grid to the ground, respectively. For auto-transformers, an effective GIC value is used (Albertson, 1981). System equivalents, i.e. inter-ties to other systems, are represented by squares.

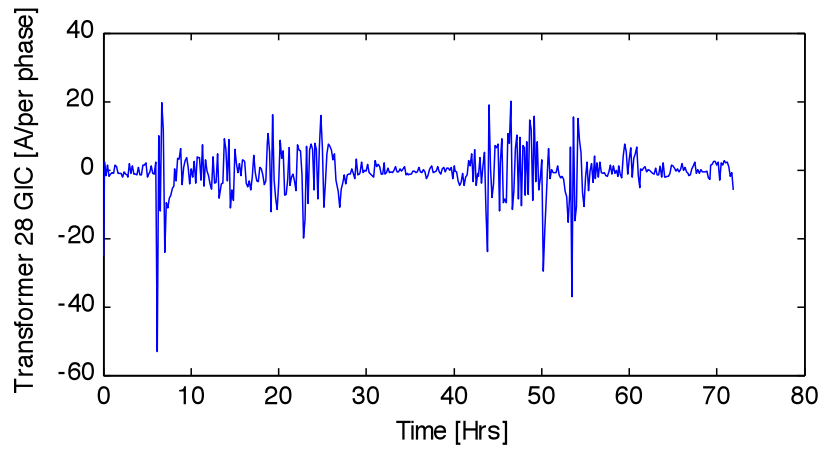


Figure 9: Time series of the modeled GIC in one of Dominion’s Virginia Power high-voltage transmission system transformers for the extreme geoelectric field scenario in Fig. 7c. The configuration of the transmission system is shown in Fig. 8. Only maximum amplitude GIC taken over 10 minute windows are shown.

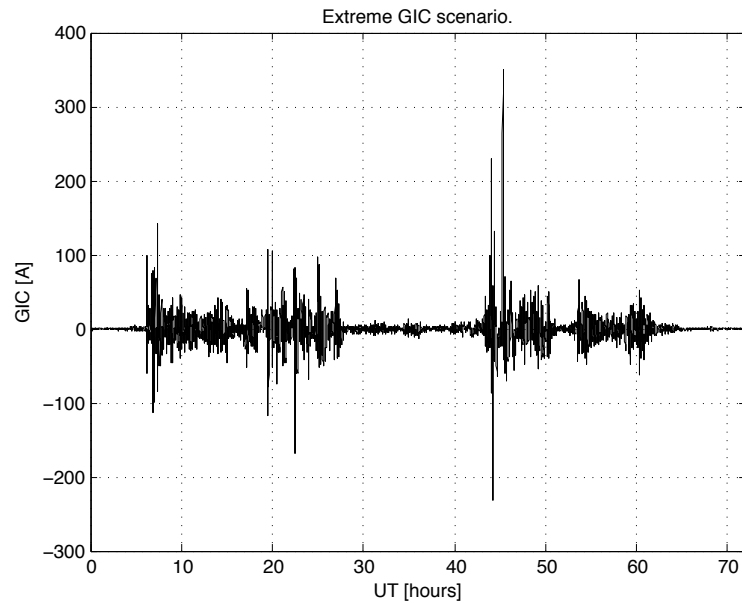


Figure 10: Extreme GIC scenario for conductive ground at location above the threshold geomagnetic latitude. System parameters  $a = b = 50 \text{ A}\cdot\text{km}/\text{V}$  were used to map the geoelectric field into GIC. See the text for details.

## 5 Discussion

In this paper we explored 100-year extreme geoelectric field scenarios by taking into account the key geophysical factors associated with the geomagnetic induction process. More specifically, we derived explicit geoelectric field temporal profiles as a function of ground conductivity structures and geomagnetic latitudes. We also demonstrated how the extreme geoelectric field scenarios can be mapped into GIC. The computed GIC can then be used in further engineering analyses that are needed to quantify the impact on the conductor systems such as high-voltage power transmission systems of interest.

Although we hope that the work in this paper provides initial input for further engineering analyses attempting to quantify the impact of extreme geomagnetic storms on high-voltage power transmission systems, it is emphasized that further work is needed to refine and improve the generated scenarios. For example, due to the poor knowledge of global spatial characteristics of the geoelectric field during extreme events, the scenarios presented in this paper apply only to regional scales of the order of 100-1000 km. It is of great interest to expand the scenarios for application in direct global calculations.

Also, one-dimensional representative ground conductivity models were used to account for varying electromagnetic response of different local geological structures. In principle, if the local ground conductivity is known, the statistics used in this paper can be tailored for specific regions providing thus significant refinement for extreme geoelectric field scenarios. It is also of interest to investigate the impact of steep horizontal ground conductivity gradients on the extreme geoelectric field scenarios. For this one would need to utilize more complex mathematical framework allowing two- or three-dimensional ground conductivity structures in the calculation of the geoelectric field. We, however, emphasize that effective one-dimensional ground conductivity models applied with the plane wave method are reasonable in many, if not in most, situations.

Perhaps the most critical and still somewhat open question that needs further clarification concerns the dynamics and the location of the identified threshold geomagnetic latitude. Since the geoelectric field amplitudes experience significant drop across the threshold, the location has significant implications for the extension of global impacts of extreme storms. Detailed studies of the historical records of extreme geomagnetic storms and accom-

panied auroral sightings as well as physics-based magnetosphere-ionosphere models capturing the key physical processes associated with the low-latitude auroral boundary may be used to shed further light on the topic.

Ultimately, only the high temporal resolution global geomagnetic recordings for very extended time periods are able to provide definitive quantification of likelihoods and spatiotemporal characterization of extreme geomagnetic storm events. The modern 60-s digital recordings are available only since about 1980s, which presents obvious difficulties in trying to extract information about 100-year events today. Regular refinement of the statistics derived in this paper should thus be carried out and over time we will be able to quantify more definitively the severity of 100-year geomagnetic storms.

Since the major of the goal of the work at hand is to facilitate further engineering analyses quantifying the extreme geomagnetic storm impact on high-voltage transmission systems, the generated extreme geoelectric field scenarios in Fig. 7 are publicly available. The digital data can be requested from A. Pulkkinen (antti.a.pulkkinen@nasa.gov).

## Appendix A: Mapping the ground geomagnetic field into the geoelectric field

Because of the central role of the process in most GIC modeling, we give here a brief overview of application of the plane wave method to mapping of the ground geomagnetic field fluctuations into the geoelectric field. The method was first formulated by Cagniard (1953) and has since been used extensively in general geomagnetic induction and magnetotelluric studies. The method is based on the concept of surface impedance, which is defined as

$$\tilde{Z} = \mu_0 \frac{\tilde{E}_x}{\tilde{B}_y} = -\mu_0 \frac{\tilde{E}_y}{\tilde{B}_x} \quad (5)$$

where  $(\tilde{E}_x, \tilde{E}_y)$  and  $(\tilde{B}_x, \tilde{B}_y)$  are the horizontal components of the electric field and magnetic field, respectively,  $\mu_0$  is the vacuum permeability and tilde indicates quantities in the spectral domain. In geophysical applications of Eq. (5) the fields are evaluated at the surface of the Earth. By assuming quasi-static temporal fluctuations, i.e. neglecting the displacement current in Maxwell's equations, and by assuming that the horizontal field gradients

vanish, the impedance of the layer  $n$  of one-dimensional layered ground can be computed using recursive formula

$$\tilde{Z}_n = \frac{i\omega\mu_0}{\gamma_n} \coth(\gamma_n d_n + \coth^{-1}(\frac{\gamma_n}{i\omega\mu_0} \tilde{Z}_{n+1})) \quad (6)$$

where  $\gamma_n^2 = i\omega\mu_0\sigma_n$ ,  $d_n$  the thickness of the layer  $n$ ,  $\omega$  the angular frequency of the field fluctuations and  $\sigma_n$  is the conductivity of the layer  $n$ . To obtain the surface impedance in Eq. (5), one sets  $n = 1$  in Eq. (6) and computes the impedance values recursively starting from the bottom of the modeled ground structure. The bottom layer is assumed infinitely thick.

While quasi-stationary approximation is valid for geomagnetic induction studies having typically frequencies of temporal field fluctuations below 1 Hz (e.g., Weaver, 1994), the assumption about vanishing horizontal gradients of field fluctuations may seem at first invalid especially during strong geomagnetic storm conditions. However, Dmitriev and Berdichevsky (1979) showed that the above formulation holds also if the plane wave requirement is relaxed into assumption about locally ( $\sim 100$  km) linear variation of the surface magnetic field. The extended validity of the “plane wave” formulation is likely one key reason for the success of the method in GIC applications.

The process for mapping the ground geomagnetic field into the geoelectric field is then as follows:

1. Convert the horizontal ground geomagnetic field into the spectral domain by using the Fourier transform.
2. Compute the surface impedance using Eq. (6).
3. Compute the spectral domain horizontal geoelectric field using Eq. (5).
4. Convert the spectral domain geoelectric field into the time domain using the inverse Fourier transform.

It is also noted that it follows from the basic properties of the Fourier transform:

$$\tilde{B} = \frac{1}{i\omega} \frac{d\tilde{B}}{dt}(\omega) \quad (7)$$

Consequently, computation of the geoelectric field from both the ground geomagnetic field and the time derivative of the geomagnetic field using Eqs. (5) and (6) is a linear operation.

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